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Hualgayoc mining district, northern Peru: Testing the use of zircon composition in exploration for porphyry-type deposits

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ABSTRACT

The Hualgayoc district in the Cajamarca region, northern Peru, hosts several deposits, including the Cerro Corona Au + Cu porphyry deposit (0.55 Mt Cu, 99 t Cu, Au) and the Tantahuatay high-sulfidation epithermal Cu deposit (74 t Cu), plus other Cu, Cu,

1. Introduction

Porphyry-style Cu \pm Au and high-sulfidation epithermal Au deposits are important sources of Cu and Au in our society. These deposits are typically associated with oxidized and hydrous arc magmas of intermediate to felsic composition (Ishihara, 1977; Streck and Dilles, 1998; Richards et al., 2012). The oxidation conditions of the metal-fertile magmas are considered to have fO_2 greater than one logarithmic unit above the fayalite-magnetite-quartz (FMQ) buffer (Wang et al., 2014; Shen et al., 2015; Hattori, 2018). This apparent association of mineralization with relatively oxidized magmas is attributed to the high capacity of oxidized magmas to extract metals and S from their magma source regions and to transport these elements to shallow crustal levels without immiscible separation of sulfides that may sequester metals and prevent partitioning to a magmatic-hydrothermal fluid (Botcharnikov et al., 2011; Zajacz et al., 2012).

Rare earth elements (REE) mostly have a valence charge of +3, but Eu is also 2+ in reduced magmas and Ce can be +4 in oxidized magmas (Burnham and Berry, 2014; Burnham et al., 2015; Smythe and Brenan, 2015). Since Eu²⁺ is preferentially incorporated into the Ca²⁺ site of

plagioclase, a melt progressively loses Eu during plagioclase crystallization. Ce⁴⁺ is easily incorporated into the Zr⁴⁺ site of zircon whereas Eu²⁺ is not readily included into the crystal structure. Thus, the magnitude of Ce and Eu anomalies in zircon reflects magmatic redox conditions (Burnham and Berry, 2012; Smythe and Brenan, 2016). Previous studies on zircon compositions from porphyry Cu deposits show small negative anomalies of Eu, Eu/Eu*, and large positive Ce anomalies, expressed as high Ce⁺⁴/Ce⁺³ (Ballard et al., 2002; Liang et al., 2006). Ballard et al. (2002) reported zircon from intrusions associated with porphyry-Cu mineralization in the Chuquicamata-El Abra area in northern Chile have $Ce^{+4}/Ce^{+3} > 300$ and Eu/Eu^* $(=[Eu_{cn}]/[(Sm_{cn}*Gd_{cn})^{0.5}]$ where cn refers to chondrite normalized values) > 0.4. Liang et al. (2006) proposed a Ce⁺⁴/Ce⁺³ ratio of 120 in zircon separating ore-bearing and barren porphyry intrusions based on the study of the Yulong Cu belt in eastern Tibet. Shen et al. (2015) noted that intermediate to large porphyry deposits (>1.5 Mt Cu) have higher $\mbox{Ce}^{+4}/\mbox{Ce}^{+3}$ in zircon (>120) than small deposits (<1.5 Mt Cu) in the Central Asian Orogenic belt. Meng et al. (2018) also suggested Ce⁺⁴/ Ce⁺³ values for zircon from ore-bearing porphyry intrusions range between 30 and 3000 and Eu/Eu* between 0.2 and 0.8. Therefore, it has

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