

1 **Paragenesis and composition of tourmaline types along the P2 fault and McArthur River uranium**
 2 **deposit, Athabasca Basin, Canada**

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7 ABSTRACT

8 The P2 fault, a 13 km long steeply dipping reverse fault, is the main structural control of
 9 the McArthur River uranium deposit in the eastern Athabasca Basin, northern Saskatchewan,
 10 Canada. Three types of tourmaline were observed in the metasedimentary basement rocks along
 11 the P2 fault: early oxy-schorl
 12 $[(\square_{0.37}\text{Na}_{0.47}\text{Ca}_{0.16})(\text{Mg}_{0.72}\text{Fe}^{2+}_{1.30}\text{Ti}_{0.07}\text{Al}_{0.91})\text{Al}_6(\text{Si}_{5.79}\text{Al}_{0.21}\text{O}_{18})(\text{BO}_3)_3\text{OH}_3(\text{F}_{0.08}\text{OH}_{0.29}\text{O}_{0.63})$,
 13 where \square = vacancy] of magmatic origin, and hydrothermal oxy-dravite
 14 $[(\square_{0.18}\text{Na}_{0.57}\text{Ca}_{0.23}\text{K}_{0.02})(\text{Mg}_{1.93}\text{Fe}^{2+}_{0.62}\text{Ti}_{0.15}\text{Al}_{0.29})\text{Al}_6(\text{Si}_{5.93}\text{Al}_{0.07}\text{O}_{18})(\text{BO}_3)_3\text{OH}_3(\text{F}_{0.20}\text{OH}_{0.23}\text{O}_{0.57})$
 15 and magnesio-foitite
 16 $[(\square_{0.77}\text{Na}_{0.20}\text{Ca}_{0.02}\text{K}_{0.01})(\text{Mg}_{1.99}\text{Fe}^{3+}_{0.07}\text{Al}_{0.92})\text{Al}_{6.00}(\text{Si}_{6.00}\text{O}_{18})(\text{BO}_3)_3(\text{OH}_3)(\text{F}_{0.04}\text{OH}_{0.71}\text{O}_{0.25})$. Oxy-
 17 schorl formed in granitic pegmatites, a partial melt product of the metasediments during their peak
 18 metamorphism. Oxy-dravite formed from hydrothermal fluids after the peak metamorphism but
 19 before deposition of the Athabasca sandstones, whereas magnesio-foitite is a product of later, low
 20 temperature hydrothermal activity. Both oxy-schorl and oxy-dravite are coarse-grained (from 500
 21 μm , up to 1 cm), whereas magnesio-foitite occurs as radial aggregates of fine, prismatic crystals
 22 ($<15 \mu\text{m}$ in width). Magnesio-foitite crystallized together with sudoite, illite and “APS” minerals

23 (alunite-supergroup LREE-rich aluminum phosphate-sulphate minerals) along the entire studied
24 length (~ 7km) of the P2 fault and is abundant in proximity to the Zone 2 ore body of the
25 McArthur River deposit. In the ore zone, the assemblage occurs with uraninite and is partially
26 overprinted by late, remobilized uraninite and sudoite. Therefore, magnesio-foitite is likely
27 contemporaneous with the main stage of uranium mineralization. It is characterized by a high
28 vacancy in its X-site (0.70 – 0.85 *apfu*) and high Al in its Y-site (0.70 – 1.12 *apfu*), suggesting that
29 magnesio-foitite likely replaced pre-existing high Al phases, such as kaolin and sudoite. The
30 occurrence of magnesio-foitite along the entire P2 fault, in areas of mineralization and apparently
31 barren areas, suggests similar fluids travelled along the entire P2 fault, but only produced ore in
32 localized areas.

33 *Keywords: tourmaline, oxy-schorl, oxy-dravite, magnesio-foitite, unconformity-type uranium*
34 *deposits, hydrothermal ore deposits, alteration*

35 INTRODUCTION

36 The Athabasca Basin hosts world-class unconformity-type uranium deposits, including the
37 McArthur River deposit. Prevalent models for uranium mineralization invoke an oxidizing, highly
38 saline brine (25 – 35 wt% NaCl equiv.), with a marine component, generated during basin
39 development (e.g., Hoeve & Sibbald 1978, Kotzer & Kyser 1995; Alexandre *et al.* 2005, Derome
40 *et al.* 2005, Richard *et al.* 2011, Mercadier *et al.* 2012). Quartz-hosted fluid inclusions in ore
41 breccias indicate that two types of fluids were present during mineralization, i) an earlier, acidic,
42 NaCl-rich, uranium-bearing brine and ii) an evolved CaCl₂-rich brine assumed to have formed by
43 the interaction of the earlier brine with basement rocks (Derome *et al.* 2005, Richard *et al.* 2010,
44 Richard *et al.* 2011, Richard *et al.* 2012). Uranium deposition was caused by the reduction of U⁶⁺
45 to U⁴⁺ in the fluids. Proposed reductants include minerals (e.g., graphite, Fe²⁺-bearing chlorite and